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1. Introduction

Title of investigation: Spherical harmonic representation of the main
geomagnetic field for world charting and
investigations of some fundamental problems of
physics and geophysics

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Progress Report, June 1st, 1981

(E82-10203) SPHERICAL HARMONIC
REPRESENTATION OF THE MAIN GEOMAGNETIC FIELD
FOR WORLD CHARTING AND INVESTIGATIONS OF
SOME FUNDAMENTAL PROBLEMS OF PHYSICS AND
GEOPHYSICS (Institute of Geological

N82-23576

Unclas
G3/43 00203

RECEIVED

DEC 1, 1981

SIS/902.6

M-037

TYPE II

2. Techniques

Data used: Magsat Investigator B test tapes

Magsat Investigator B data tapes

Tapes of selected data on 15 magnetically quiet days

Spherical harmonic models of the geomagnetic field produced from Magsat data by GSFC

Programs used: GSFC program for analysing selected quiet day data

Analysis program producing spherical harmonic models of the geomagnetic field and including, besides the usual internal poloidal field, external and toroidal fields

Various test programs produced by IGS and Liverpool groups

3. Accomplishments

Investigator B test tapes were read successfully. A start had been made on a preliminary survey of the first batch of Investigator B data tapes when it was discovered (and later verified by NASA) that these were not error-free. It was decided to postpone further work on these tapes until corrected versions were received. These finally arrived at the end of the reporting period.

The selected quiet-day data are of particular interest to the Liverpool group. They have experienced some difficulties with the GSFC analysis program, particularly concerning the altitude correction. Tests of Maxwell's equations against other proposed modified equations have begun and initial results look promising.

The 1980 World Chart spherical model has been compared with the MGST (3/80) and MGST(6/80) models. Agreement in declination is satisfactory (to within $\pm 0.5^\circ$) between latitudes of 50°N and 50°S , the region where Magsat vector data were used in the MGST models.

4. Significant results

The investigation is at too earlier a stage for there to be any really significant results to be announced. The agreement of the 1980 World Chart model with the Magsat data (as represented by the MGST(3/80) and (6/80) models) is reassuring, however.

5. Publications

Barracclough, D.R., 1981. Geomagnetic field modelling using satellite data, *Nature*, 290, 86.

6. Problems

Delays in delivery of Investigator B data tapes and errors in the first batch of tapes have caused hold-ups in the investigation.

7. Data quality and delivery.

See remarks in Section 6.

morphologically. As mentioned above, such chemotaxonomic considerations would today raise a red flag against halogeton. Another example is the sicklepod milkvech (*Astragalus falcatus* Lam.), which was introduced around 1920 and has been recommended for range plantings for over 25 years; recognition that many *Astragalus* species contain toxic concentrations of nitro compounds led to identification of sicklepod milkvech as poisonous, and efforts are now being directed to its eradication. Fourth,

introduced plants that are suspected of being poisonous may be fed to experimental animals, as extracts, dried plants, or residues.

With current pressures to strip-mine for coal, there is much interest in finding plant species that will grow on mine spoils, reproducing rapidly and providing erosion control, and, if possible, grazing for livestock and food and cover for wildlife. The clear need for such plants to be non-poisonous gives special urgency to the programme outlined above. □

in the spherical harmonic expansion, which Backus highlighted in his uniqueness studies, are likely to be relatively poorly determined in an analysis of F data alone. These series (which Stern and Bredekamp called 'Backus series') consist of terms which all have the same value of m . Each series starts with a sectional ($n = m$) term and succeeding terms have values of n equal to $m + 2, m + 4, m + 6$, and so on. In agreement with the findings noted above, Stern and Bredekamp found, using simulated data, that the terms with $m = n$ had the largest uncertainty for a given series.

Now that data are available from Magsat, it is possible to study these effects using real data. Stern *et al.* (*Geophys. Res. Lett.* 7, 941; 1980), in one of the first reports of results from Magsat, have made an initial attempt at this. They have compared two models, one based on vector data from early in the Magsat mission (November 1979) and the other based only on scalar (F) data collected at about the same time. Values of Z computed from the two models differed by up to 2,300 nT at the Earth's surface and by up to 1,500 nT at the satellite's altitude. (For comparison, typical mid-latitude values of Z are of the order of 40,000 nT.) The differences, when plotted on a world map, showed a pattern of six foci, all positioned very near to the dip-equator. When the two models were compared coefficient by coefficient, those in the Backus series described above did indeed show larger differences than the coefficients not belonging to such series. There were suggestions that other effects were present affecting the 'non-Backus' terms and work is in progress in this area.

Thus, the doubts cast on geomagnetic field models based solely on F data seem to be confirmed. Fortunately there are relatively few such models and most of these have been produced for specific purposes related to the scalar data on which they were based. Most producers of models for navigational and general scientific purposes (for example, Barraclough *et al.* *Geophys. J. R. astr. Soc.* 43, 645; 1975; Barker *et al.* *Geophys. J. R. astr. Soc.* in the press) have been careful to use F observations as only a part of the input data and their models will therefore be free from the deficiencies discussed above. Potential users of field models should be wary in their choice of models. Field models based solely on F data can be used with some confidence if one is interested only in the intensity of the geomagnetic field. If vector information is needed it is best to avoid such models and to use instead one of the models based on component data as well as F data. The International Association of Geomagnetism and Acronomy, at its next scientific assembly this summer, plans to recommend an International Geomagnetic Reference Field which should provide a set of reliable models describing the geomagnetic field since 1965. □

Geomagnetic field modelling using satellite data

from David R. Barraclough

ACCURATE DESCRIPTIONS of the Earth's magnetic field are useful for navigational purposes and form the basis for many scientific studies. Such descriptions usually take the form of mathematical models, most often in terms of an expansion in spherical harmonics, and magnetic charts can readily be produced from the models. A major difficulty in producing geomagnetic field models has always been the acquisition of an up-to-date and well distributed set of accurate observations. As a satellite in a polar orbit can survey the entire globe in a matter of days, it was an obvious step to use such a satellite to measure the geomagnetic field.

The most recent satellite survey of the geomagnetic field was performed by the NASA/US Geological Survey satellite Magsat between November 1979 and June 1980. Magsat differed from previous geomagnetic survey satellites in that it was able to measure not only the strength of the field but also its direction.

Since the geomagnetic field is derivable from a scalar potential which satisfies Laplace's equation, it was possible to use the earlier satellite observations of F to refine existing models of the field. Although such models fitted the input F data very well (with a r.m.s. residual of about 10 nT), there was evidence that they did not describe the vector geomagnetic field as accurately (Ben'kova *et al.* *Bull. int. Ass. Geomag. Aeron.* No. 152-163; 1971; Fabiano & Peddie *J. geophys. Res.* 76, 3816; 1971). These findings led to several theoretical studies of the production of field models from F observations alone.

Hurwitz and Knapp (*J. geophys. Res.* 79, 3009; 1974) and Barraclough and Nevitt (*Phys. Earth planet. Interiors* 13, 123; 1976), using simulated data with noise added, verified that good fits to F did not

guarantee a good description of the other field components. In particular, large errors in the vertical component (Z) of the field occurred near the dip-equator (the line of zero inclination). In terms of the coefficients of the spherical harmonic model, those having equal values of m and n were least well determined. These are the sectorial harmonics, which divide the surface of a sphere into segments like those of an orange. It was also shown that the large errors could be removed by using small amounts of vector data (about 10 per cent of the total) covering the equatorial region.

Hurwitz and Knapp (*op. cit.*) and Lowes (*Geophys. J. R. astr. Soc.* 42, 637; 1975) showed that the root of the problem lay with what they termed the perpendicular error effect. When fitting only a single component of a vector field, the perpendicular components synthesized from the resulting model can have errors associated with them which are considerably larger than the errors in the component fitted. This results from the lack of constraints in the perpendicular directions. In some cases, for example the production of models from Z data only, the perpendicular error effect is of minor importance because of the particular configuration of the geomagnetic field. It turns out, however, that fitting F data is a particularly bad choice. The perpendicular error effect leads to large errors in Z and these are especially troublesome where Z is small, that is, near the dip-equator, as noted above.

Backus (*Q. Jl Mech. appl. Math.* 21, 195; 1968; *J. geophys. Res.* 75, 6339; 1970; *Geophys. Res. Lett.* 1, 21; 1974) investigated whether models derived from F data alone were unique and showed that, under certain conditions, they were not. In practical field modelling using real data these conditions for non-uniqueness are not fulfilled. Stern and Bredekamp (*J. geophys. Res.* 80, 1776; 1975) have, however, shown that certain series of terms

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